

Lectures 3: Evolution of metabolism, the atmosphere.
Reading: Chapter 2 and 3 of Schlesinger

1. Origin of organic molecules
 - a. Exploitation of surface element reservoirs (general focus on most abundant elements).
 - b. Special “polymer” chemistry of carbon, information, functionality optimization of complex C-C chains and rings.
 - c. Carbon chemistry supplemented by elements with similar & higher electronegativity (N, H, O) – chain-compatible elements
2. Origin of metabolism
 - a. All metabolic activity requires input of energy
 - i. Photosynthesis – E from light
 - ii. Chemosynthesis – E from chemical disequilibrium in environment
 - iii. Heterotrophy – E from other organisms
 - b. Early Earth probably had small reservoirs of free organic molecules (e.g., CH₃COOH, amino acids), from either extraterrestrial sources (meteorites, comets) or terrestrial sources (lightning, hydrothermal reactions, etc.)
 - c. Atmosphere & oceans lacked free O₂, increasing lifetimes of organic molecules.
 - i. Spontaneous exploitable disequilibria
 1. CH₃COOH → CO₂ + CH₄ (Methanogenesis 1)
 - d. Initial chemosynthesis also possibly based on spontaneous disequilibria
 1. CO₂ + 4H₂ → CH₄ + 2H₂O (Methanogenesis 2)
 - ii. Modern habitats?
 - e. Photosynthesis – harvesting energy from sunlight, and storing it by *molecular separation* of oxidants and reducers.
 - i. Oxidant: substance that tends to react by taking electrons.
$$\text{O}^0_2 + 4\text{e}^- \rightarrow 2\text{O}^{2-}$$
 - ii. Reducers: substance that tends to react by giving up electrons.
$$\text{H}^0_2 \rightarrow 2\text{H}^+ + 2\text{e}^-$$
 - iii. Combining an oxidant with a reducer liberates energy.
Conversely, energy is required to separate them.
$$2\text{H}_2 + \text{O}_2 \rightarrow \text{H}^{(+)}_2\text{O}^{(2-)}$$

Quantifying energy (thermodynamics) Reactions “go” spontaneously if their free energy is negative. I.e., the total free energy of the products (right side) must be lower than the total free energy of the reactants (left side) (See HW).

Ch. 3: The atmosphere.

1. The Earth's atmosphere is the least massive and most rapidly mixed global geochemical reservoir.
 - a. Mass $5.148 \pm 0.001 \times 10^{18}$ kg, variable due to water vapor
 - b. Mix time of lower atmosphere ~ 1 year (upper atmosphere slower)
 - i. – little geographic variation in composition (except H₂O)
 - c. Major constituents (ignoring H₂O vapor)
 - i. N₂ – 78%
 - ii. O₂ – 21%
 - iii. Ar – 0.9%
 - iv. CO₂ -- .039% (rising $\sim 0.0002\%$ /year)
 - v. H₂O – variable, typically $\sim 1\%$ at ground level
 - vi. Trace gasses – CH₄, H₂, N₂O, O₃, Ne, He, Kr, Xe, etc.

Average composition of dry atmosphere, by volume	
Gas	per NASA
Nitrogen , N ₂	78.084%
Oxygen , O ₂	20.946%
Argon , Ar	0.934%
Minor constituents (in ppm)	
Carbon Dioxide , CO ₂	383
Neon , Ne	18.18
Helium , He	5.24
Methane , CH ₄	1.7
Krypton , Kr	1.14
Hydrogen , H ₂	0.55
Water	
Water vapour	Highly variable; typically makes up about 1%

2. Basic chemistry of air – nearly ideal gas, $PV \approx nRT$
 - a. Expands on heating ($V \approx T$ at constant pressure)
 - b. Temp increases when air is compressed, drops when expanded
 $T_2/T_1 \approx (V_1/V_2)^{2/5}$ (dry, adiabatic – no heat added)
 - c. Capable of holding more H₂O at high T
 - i. Warming or descending air evaporates water
 - ii. Cooling or ascending air condenses water (clouds)
3. Structure of the atmosphere (Fig. 3.1)
 - a. Troposphere \sim adiabatic (warm low, cold high)
 - i. Lowest ~ 10 km of atm.
 - ii. 80% of atmospheric mass