#### Announcements

- •Reading: p.155-160
- •Important not to fall behind in reading or assignments!
- •Turn in field trip notes at the end of next week
- •Gimme your pictures
- •Homework 3 due May 8

Harker diagram for Crater Lake

Closure problem!

What does this mean?

Figure 8-2. Harker variation diagram for 310 analyzed volcanic rocks from Crater Lake (Mt. Mazama), Oregon Cascades. Data compiled by Rick Conrey (personal communication).



### **Pearce Diagrams**

- Molar units (elements or oxides)
- Ratio to a "perfectly" incompatible element (stays in melt)
- Way to test a hypothesis, or negate one
- Not a way to prove fractional crystallization



**Figure 8–4** Pearce element diagram of 0.5(Fe + Mg)/K vs. Si/K for two Hawaiian picritic magma suites. From Nicholls and Russell (1990).

#### The Basalt Tetrahedron and the Ne-Ol-Q base





**Figure 8-12.** Left: the basalt tetrahedron (after Yoder and Tilley, 1962). J. Pet., 3, 342-532. Right: the base of the basalt tetrahedron using cation normative minerals, with the compositions of subalkaline rocks (black) and alkaline rocks (gray) from Figure 8-11, projected from Cpx. After Irvine and Baragar (1971). Can. J. Earth Sci., 8, 523-548.



Thermal divide separates the silicasaturated (subalkaline) from the silicaundersaturated (alkaline) fields at low pressure

Cannot cross this divide by FX, so can't derive one series from the other (at least via low-P FX)



#### Ionic radii

CrystalMaker Element Tables Shannon & Prewitt "Cr										stal"	Radi	ii		http://v	www.cry	vstalmaker.com
H H				F	e 1	.25	Å									He
3 4 Li Be ♥				F	e <sup>2+</sup>	0.7	7 Å				5 B	6 C	7 N	8 0	9 F	10 Ne
11 12 Na Mg				F	e <sup>3+</sup>	0.6	3 Å				13 Al •	14 Si	15 P	16 S	17 CI	18 Ar
19 20 K Ca	21 Sc	22 Ti	23 V	24 Cr	25 Mn	26 Fe	27 Co	28 Ni	29 Cu	30 Zn	31 Ga	32 Ge	33 As	34 Se 🎐	35 Br	36 Kr
37 38 Rb Sr	39 Y	40 Zr	41 Nb	42 Mo	43 Tc	44 Ru	45 Rh	46 Pd	47 Ag	48 Cd	49 In 🎱	50 Sn	51 Sb	52 Te	53 I	54 Xe
55 56 Cs 56 Ba	71 Lu	72 Hf	73 Ta	74 W	75 Re	76 Os	77 Ir 🍑	78 Pt	79 Au	80 Hg	81 TI	82 Pb	83 Bi	84 Po	85 At 🎱	86 Rn
Image generated by CrystalMaker 6.3:	57 La	58 Ce	59 Pr	60 Nd	61 Pm	62 Sm	63 Eu	64 Gd	65 Tb	66 Dy	67 Ho	68 Er	69 Tm	70 Yb	:	Scale 10 Å
nteractive crystal & molecular structures visualization for Mac OS X	89 Ac	90 Th	91 Pa	92 U	93 Np	94 Pu	95 Am	96 Cm	97 Bk	98 Cf	99 Es	100 Fm	101 Md	102 No		

 Cations are smaller than neutral atoms, anions are larger. For a given element, size decreases with increasing charge.

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interactive crystal & molecular structures visualization for Mac OS X	89 Ac	90 Th	91 Pa	92 U	93 Np	94 Pu	95 Am	96 Cm	97 Bk	98 Cf	99 Es	100 Fm	101 Md	102 No				

 Ionic radius increases down a chemical group. Many exceptions: REE's, many d-block transition elements

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 For ions with same electronic structure within a period (i.e., Na<sup>+</sup>, Mg<sup>2+</sup>) size decreases with increasing charge (increasing attractive charge of nucleus)



- Basic mineral structures can be understood by looking at the relative sizes and the charges of different elements in their common oxidation states!
- Nobel Prize in Chemistry 1954:
  - "for his research into the nature of the chemical bond and its application to the elucidation of the structure of complex substances"
- Nobel Peace Prize 1962



#### 1. The Coordination Principle

- Cation-anion distance is determined by the sum of the cation and anion radii
- Number of anions coordinating with cation is determined by relative size of cation and anions.
  - Cations want to fit snugly in anions- equal bonding environments for all.
  - Assumptions!

RR = Rc/Ra



UA Downs

# 2. The Electrostatic Valency Principle

Total strength of valency bonds that reach an anion from all neighboring cations = charge of anion.



#### 3. Sharing of Polyhedral Elements I

The stability of ionic structures is decreased when coordination polyhedra share edges and faces.

Melting temperature of

Forsterite = 1890 °C

Enstatite = 1550 ° C

#### 4. Sharing of Polyhedral Elements II

In a crystal containing different cations, large valence, small CN cations tend not to share polyhedral elements with each other.

More than half of ionic charge is occupied with small, highly-charged cation:  $CO_3^{2-}$ ,  $PO_4^{3-}$ 

#### 5. Principle of Parsimony

Nature tends towards simplicity!

Mineral structures tend to be simple, with only one or a few different cation sites – but each may be accommodate many different cations.

M1 = Mg, Fe<sup>2+</sup>, Fe<sup>3+</sup>, Al M2 = Mg, Fe<sup>2+</sup>, Ca, Na, Li



#### Goldschmidt's Rules: general guidelines





Partition coefficient:
 K<sub>D</sub> or D (mineral-melt)

 $K_D^A = [A]_{mineral}/[A]_{melt}$ 

$$K_D^B = [B]_{mineral}/[B]_{melt}$$



### Example: olivine

- Na : 0.02
- Mg : 6.1
- Ni : 14

 Expect to see a Na olivine? Ni?

- Partition coefficient:
  K<sub>D</sub> or D (mineral-melt)
- $K_D^A = [A]_{mineral}/[A]_{melt}$

$$K_D^B = [B]_{mineral}/[B]_{melt}$$